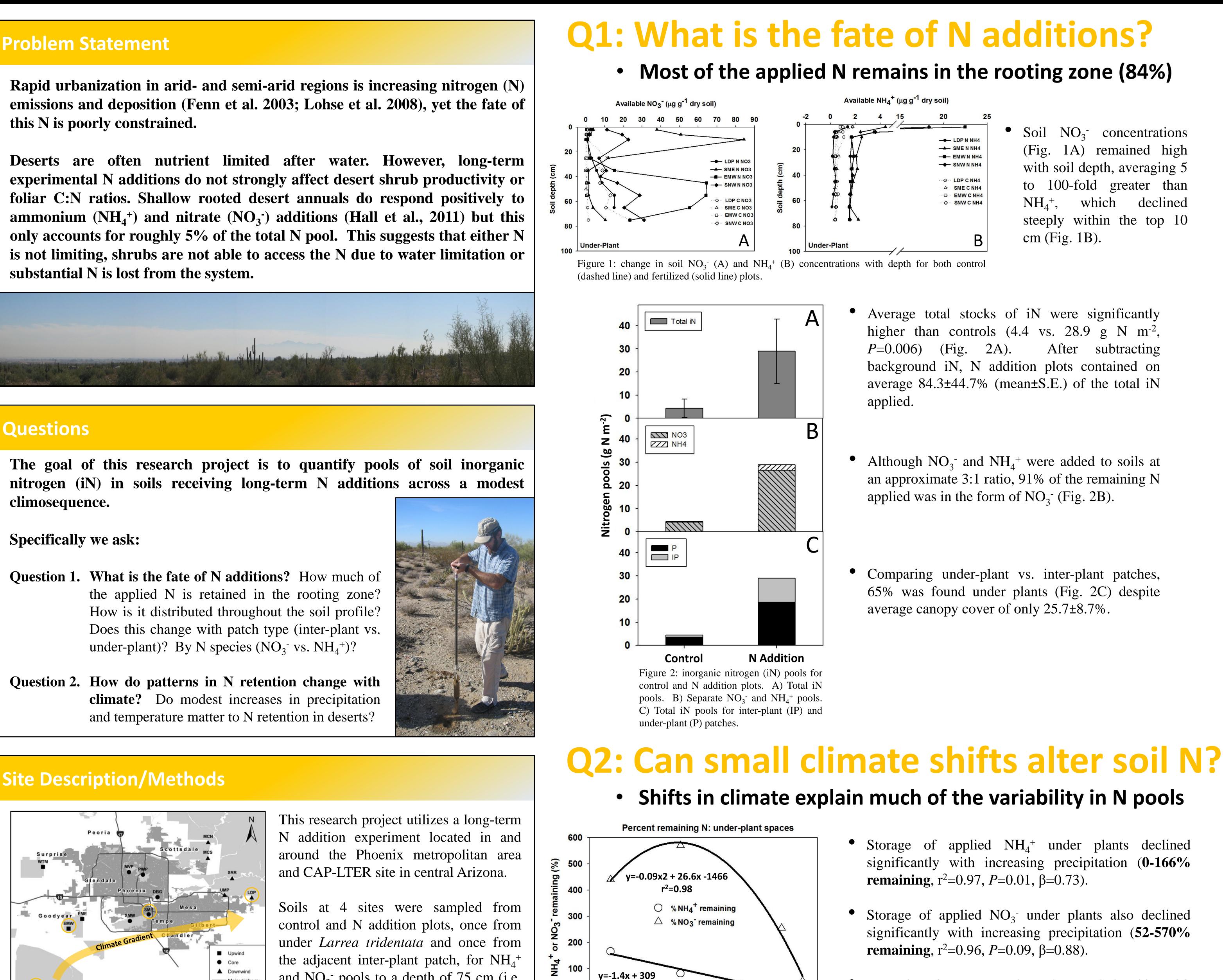


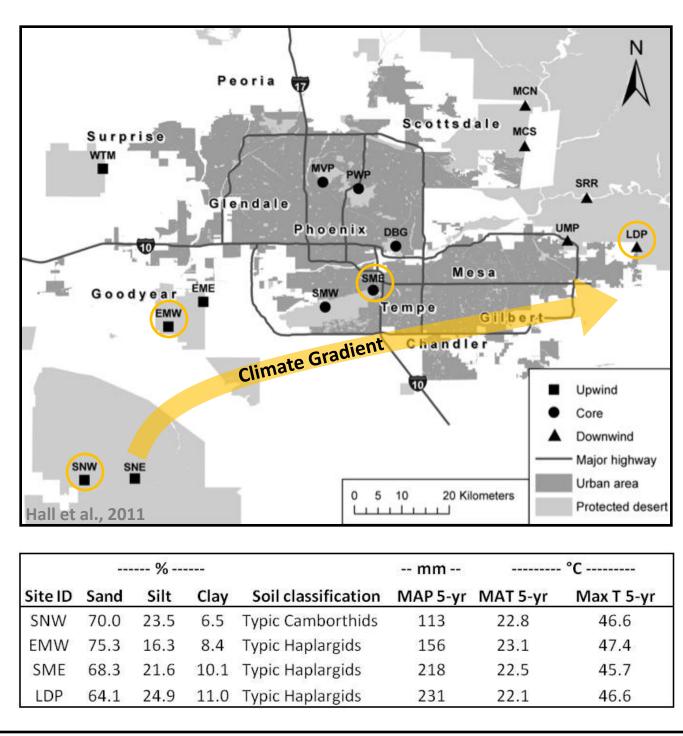
Climate Controls the Fate of Anthropogenic Nitrogen Additions in Desert Ecosystems



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and NO_3^- pools to a depth of 75 cm (i.e. rooting zone) at intervals of 2-10 cm.

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Figure 3: percent applied NO_3^- (triangles) and NH_4^+ (circles) remaining as a function of mean annual precipitation (MAP) across the climosequence.

5-yr average MAP (mm)

r²=0.97

Soil NO₃⁻ concentrations (Fig. 1A) remained high with soil depth, averaging 5 to 100-fold greater than NH_4^+ , which declined steeply within the top 10 cm (Fig. 1B).

Average total stocks of iN were significantly higher than controls (4.4 vs. 28.9 g N m⁻², P=0.006) (Fig. 2A). After subtracting background iN, N addition plots contained on average 84.3±44.7% (mean±S.E.) of the total iN

Although NO_3^- and NH_4^+ were added to soils at an approximate 3:1 ratio, 91% of the remaining N applied was in the form of NO_3^- (Fig. 2B).

Comparing under-plant vs. inter-plant patches, 65% was found under plants (Fig. 2C) despite average canopy cover of only 25.7±8.7%.

Storage of applied NH_4^+ under plants declined significantly with increasing precipitation (0-166%) **remaining**, $r^2=0.97$, P=0.01, $\beta=0.73$).

Storage of applied NO_3^- under plants also declined significantly with increasing precipitation (52-570%) **remaining**, $r^2=0.96$, P=0.09, $\beta=0.88$).

Inter-plant iN storage showed no relationship with MAP but average maximum daily temperature during summer trended strongly with inter-plant soil $NO_3^$ pools ($r^2=0.80$, data not shown).

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Site	Patch	Applied NO ₃ -	Me
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552 302 012 012 010 010 010 010	IP	10.9	
SNW	Р	2.1	
	IP	24.8	
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Conclusions

- changes in MAP (70 mm) between sites.

Nitrogen loss processes sensitive to soil moisture and temperature and likely responsible for variation in N pools between sites include: fluxes of ammonia, nitric oxide and nitrous oxide (Hall et al., 2008; McCalley and Sparks, 2008). Nitrate leaching below the rooting zone at the wetter sites may also be a mechanism of N loss (Walvoord et al., 2003).

References/Acknowledgements

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esses explain enrichment of elative to N additions?

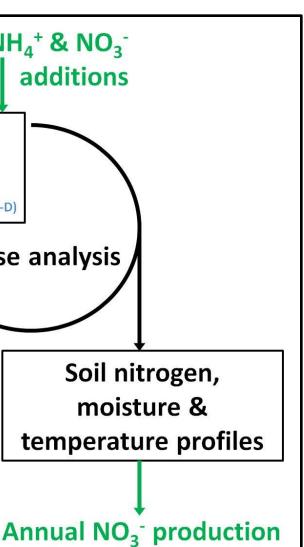


Figure 4: soil hydrologic properties were estimated from soil texture bulk density measurements pedo-transfer function. Parameter estimates from the NRCS database were used to constrain the model. Inverse model results were compared to NRCS values and $NO_3^$ profiles for hydrologic water potential over estimate soil the 5 year study. These results were incorporated into a kinetic nitrification model, which scales potential nitrification rates by soil water potential and temperature.

ologic model (Hydrus 1-D) with a simple kinetic adapted from Stark and Firestone (1996) to NO₃⁻ production under predicted soil-water e 4).

	g NO ₃ - m ⁻² 5yr ⁻¹	%
asured	NO ₃ -	Excess
NO ₃ -	production	explained
5.4	2.7	242
1.0	2.8	1356
9.3	11.8	149
13.1	7.6	248

Table 2: pools of applied and measured NO_3^- , NO_3^- produced due to nitrification during the 5 year experiment and percent of excess NO_3^- explained by nitrification model at the wet (LDP) and dry (SNW) N addition sites.

tes of nitrification rates in the upper 75 cm ssive NO_3^- observed (Table 2).

• We show that after 5 years of experimental N additions, applied N largely remain within the rooting zone (84%) of these desert soils.

• However, NH_4^+ and NO_3^- pools are strongly controlled by modest

• Our modeling results suggest that nitrification largely explains the presence of excess NO_3^{-} in N addition plots. In addition, our modeling suggests some of the NH_4^+ that was deposited in interplant spaces must have been redistributed to under-plant patches, supporting the conceptual model introduced by Hall et al. (2011).